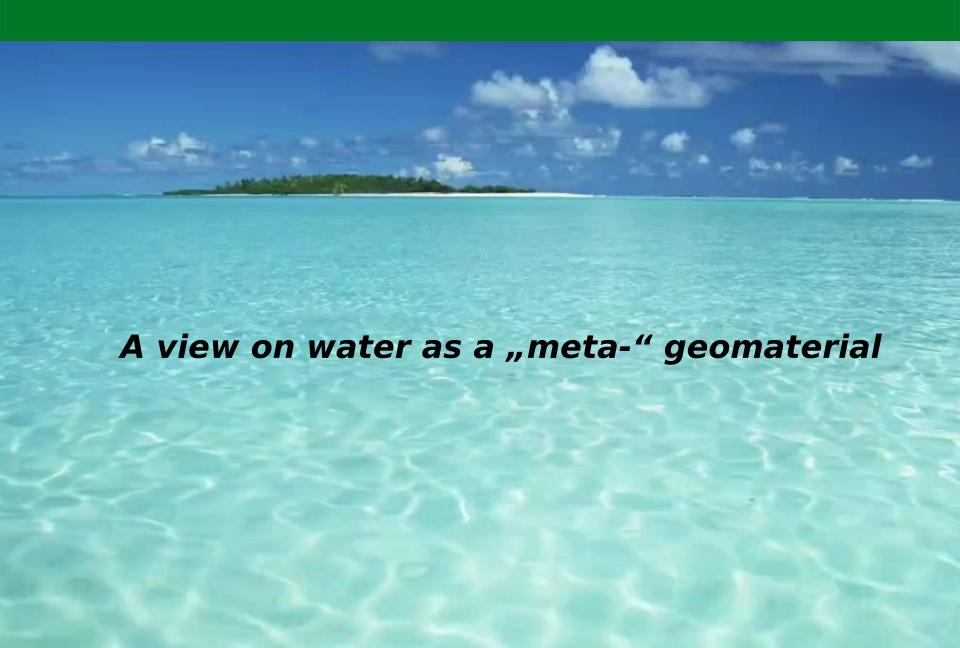
Water as "meta"-geomaterial (F. Trixler)



Water as "meta"-geomaterial (F. Trixler)



- 1. The origin of water
- 2. How was water delivered to Earth?
- 3. What let it remain on Earth?
- 4. Mineral evolution and the role of Water
- 5. Importance of water for geomaterials
- 6. Co-evolution of geo- and biosphere driven by water

1. The origin of water





Interstellar Medium (ISM)

- Mainly H and He
- 0,1 % C, N, O, Mg, ... Fe
 - → from stellar nucleosynthesis
 - → Injected into the ISM via stellar outbursts, stellar winds, red supergiant stars
 - → crystallization from gas phase (e.g. outer atmosphere of Red Giants) generates refractory component of nm/ μm size **dust particles:** Silicates (Olivine, Pyroxene), Carbides, Diamond, Graphite, PAHs...



Lynd 1251 (c) A. Jerahian

Interstellar Medium (ISM)



Lynd 1251 (c) A. Ierahian

Density of dust clouds can be so high that it is opaque for most electromagnetic radiation of nearby stars

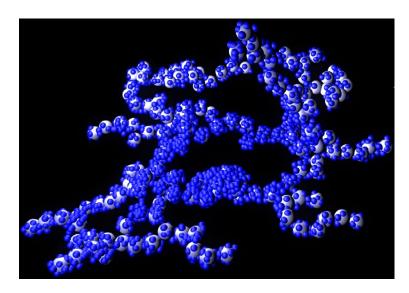
- \rightarrow low temperature (<30 K): **inefficient** gas phase synthesis
- → rich chemistry takes place; **high abundance** of H₂, H₂O, ...

} paradox

→ surface of interstellar <u>dust grains</u> catalyzes H₂ / H₂O formation



- H adsorption → exchange energy with surface → diffusion →
- → combination with other adsorbed H → H2 ... (+ OH → H2O + H)



Phys. Rev. Lett. 124, 221103 (2020).

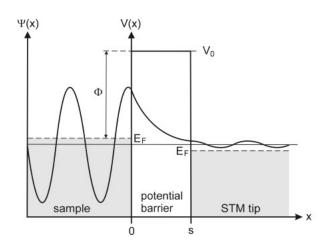
Dark Clouds in ISM

- → Low temperature: T < 20 K
- Physisorption (without fast desorption)
- But low mobility for diffusion!
- However: high efficiency in H2 production

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♦ Quantum Tunneling



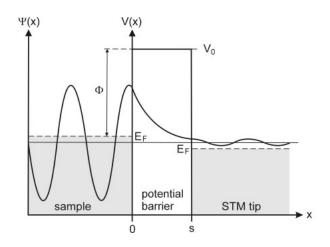
Curr. Org. Chem. 17 (2013), 1758

- → Icy mantles around interstellar dust particles (T < 20 K)
- Most abundant: H2O
- Considered as most common surface reaction:
 OH + H₂ → H₂O + H
- Large activation barrier: at T> 20 K
 - → reaction via thermal activation: very low probability

→ Icy mantles around interstellar dust particles (T < 20 K)</p>

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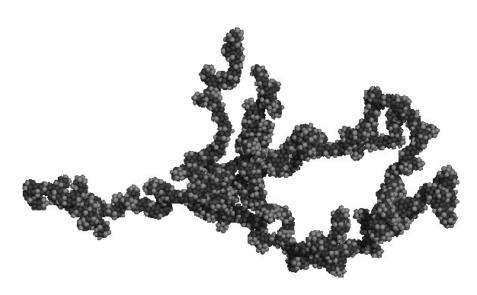
Quantum Tunneling



Curr. Org. Chem. 17 (2013), 1758



origin of water is linked to mineral particles

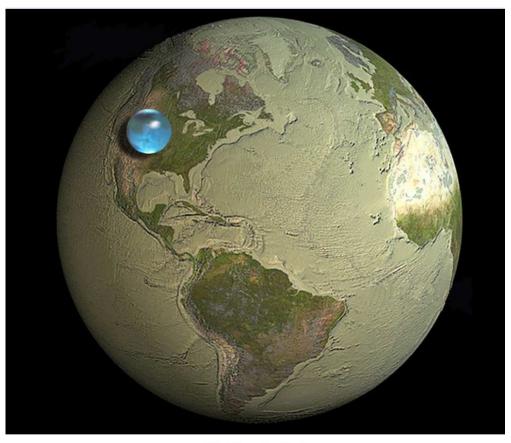


chm.bris.ac.uk

2. How was water delivered to Earth?

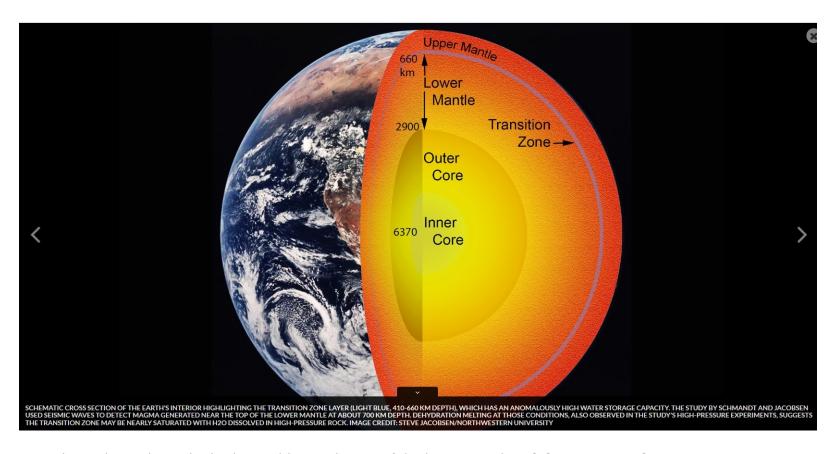


Amount of water on/near surface



All the Water on Planet Earth
Illustration Credit & Copyright: Jack Cook, Woods Hole Oceanographic Institution, Howard Perlman, USGS

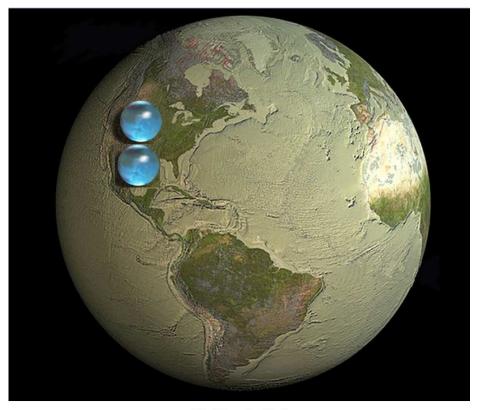
Hydration of mantle transition zone: 1 – 1.5 w% H₂O \rightarrow in sum may \sim equal oceanic water



B. Schmandt, et al., "Dehydration melting at the top of the lower mantle," **doi: 10.1126/science.1253358** D. G. Pearson, et al., "Hydrous mantle transition zone indicated by ringwoodite included within diamond," **doi:10.1038/nature13080**

Total amount of water: $\sim 0.02* (0.04 \%)$ of total mass

* M. Willians, Universe Today (2014), 12



All the Water on Planet Earth

Illustration Credit & Copyright: Jack Cook, Woods Hole Oceanographic Institution, Howard Perlman, USGS

Snow line

(ice line, frost line)

- Is the distance to the sun where T °C low: volatiles condense (in space)
- Current snow line water: ~ 2.7 AU (Ceres: 2,77 AU)
 - → Snow Line separates inner solar systems (rocky plantes) and outer solar systems (Gas Giants, Icy Moons, Ceres)

w% Water:

Earth: 00.04

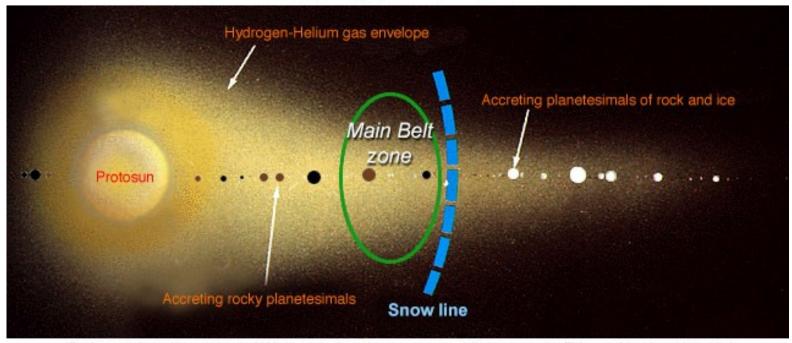
Titan: 40.00

Rhea: 75.00

Snow line

(ice line, frost line)

Solar System Formation



(Background graphic courtesy of Windows to the Universe, www.windows.ucar.edu. Ellipse added for emphasis.) psrd.hawaii.edu/WebImg/protoplanets-snowline.jpg

snow line → visible



Snow line

(ice line, frost line)



"Geomaterials in technology and environment" would be very different on an icy planetary body in the outer solar system.

Surface of Titan (Saturn System):

- water ice rocks
- sand pebbles made of icy grains and Tholins
- pores of wet sand filled with liquid methane

Nature 438 (2005), 756



availability/ abundance of certain geomaterials on the surface also depends on the location relative To the snow line of water



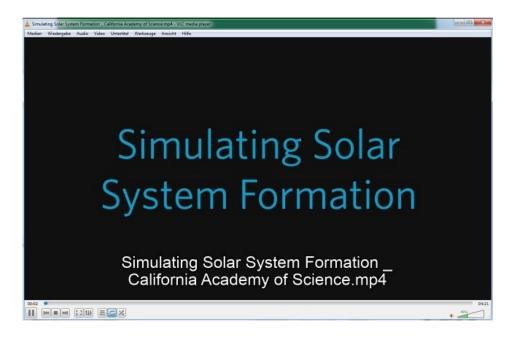
- Water abundant in the outer solar system beyond snow line
- Carbonaceous chondrites highest content of volatiles such as water
- aggregate in comets, asteroides, planetesimals
- Isotope geochemistry:
 - → Mo:

Earth accreted carbonaceous bodies late in its growth history, probably through the Moon-forming impact [http://dx.doi.org/10.5772/50172]

→ Se:

signature of Se in Earths mantle matches nearly perfect with Se signature of carbonaceous chondrites [https://dx.doi.org/10.1038/s41561-019-0414-7]

Most of water on Earth delivered by late accretion of carbonaceous chondrites from the outer solar system

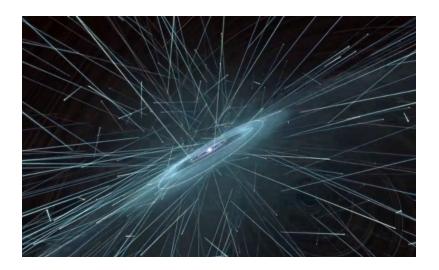


youtube.com/watch?v=yXq1i3HlumA



Water mainly delivered from outer solar system

- → packing: icy mantles on dust particles, aggregation
- → transport: C-type asteroides/planetesimals (and a bit from comets)
- → delivery service: Jupiter and other gas giants



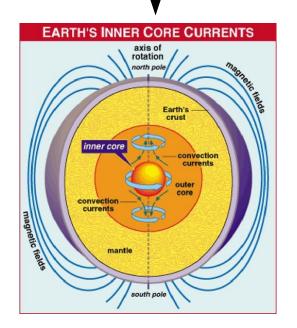
What let water remain on Earth?



• Abundance of elements critical to geophysical development:

→ U, Th, Fe, H₂O

 Determine e.g. gravity, core size, internal temperature, plate tectonics and magnetic field





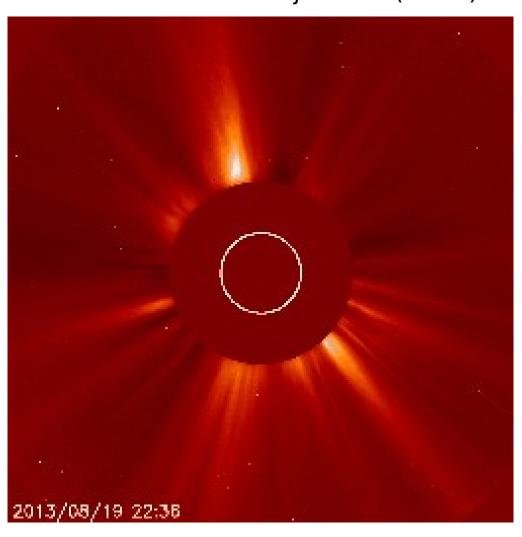
Comparing Earth / Venus

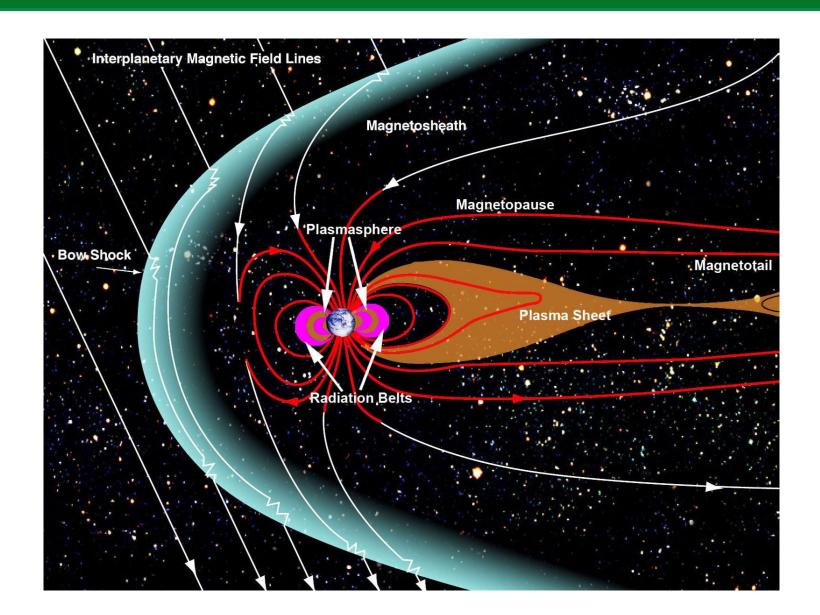
- Earth models reach steady state carbon cycles after several hundred million years of degassing, weathering, and subduction.
- Faster **plate tectonic** speeds and erosion rates tend to dampen fluctuations in the carbon cycle via silicate weathering.
- Venus models evolve to runaway greenhouses with all volatiles in the atmosphere and rapid water loss to space.
- A strong magnetic field at Venus would have prevented significant water loss.

P.Driscol, ID.Bercovici: Divergent evolution of Earth and Venus: Influence of degassing, tectonics, and magnetic fields

doi.org/10.1016/j.icarus.2013.07.025

Coronal mass ejections (CMEs)



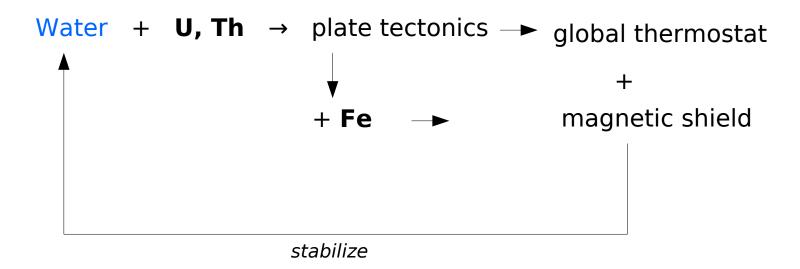


No plate tectonics → no oceanic ridges heat loss → planetary core stays hot inhibits development of magnetic field

No magnetic field → no shielding against solar wind → atmospheric loss

Water instable when:

- Atmospheric loss (no magnetic shielding)
- Runaway greenhous effect (no negative feedback between atmospheric temperature, carbon dioxide weathering and surface tectonics)



Indirect, but strong relationship between the geomaterials

U, Th and Fe, Ni with water

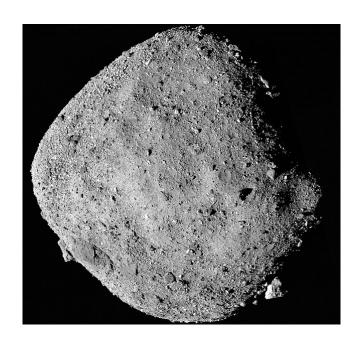
→ stabilize water on Earth via feedback loops (carbon cycle/ plate tectonics) and magnetic shielding





4. Mineral Evolution and the Role of Water

Mineral Evolution and the Role of Water



Asteroid "Bennu": ~ 250 mineral species (CI carbonaceous chondrite)



Earth: $> \sim 4.300$ mineral species

4. Mineral Evolution and the Role of Water

Hazen et al.: American Mineralogist 93, 1693–1720, 2008.

Stage	Age (Ga)	Examples of minerals ~ Cumul	ative no. species
	The era of p	planetary accretion (>4.55 Ga)	
1. Primary chondrite minerals	>4.56 Ga	Mg-olivine/pyroxene, Fe-Ni metal, FeS, CAIs	60
2. Planetesimal alteration/differentiation	>4.56 to 4.55 Ga		250
a) aqueous alteration		phyllosilicates, hydroxides, sulfates, carbonates, halite	
b) thermal alteration		albite, feldspathoids, biopyriboles	
c) shock phases		ringwoodite, majorite, akimotoite, wadsleyite	
d) achondrites		quartz, K-feldspar, titanite, zircon	
e) iron meteorites		many transition metal sulfides and phosphates	
	The era of crust ar	nd mantle reworking (4.55 to 2.5 Ga)	
3. Igneous rock evolution	4.55 to 4.0 Ga	5	350 to 500
a) fractionation		feldspathoids, biopyriboles (volatile-poor planets)	350
b) volcanism, outgassing, surface hydration		hydroxides, clay minerals (volatile-rich planets)	500
4. Granite formation	4.0 to 3.5 Ga		1000
a) granitoids		quartz, alkali feldspar (perthite), hornblende, micas, zircon	
b) pegmatites		beryl, tourmaline, spodumene, pollucite, many others	
5. Plate tectonics	>> 3.0 Ga		1500
a) hydrothermal ores		sulfides, selenides, arsenides, antimonides, tellurides, sulfosalts	
b) metamorphic minerals		kyanite, sillimanite, cordierite, chloritoid, jadeite, staurolite	
6. Anoxic biological world	3.9 to 2.5 Ga		1500
a) metal precipitates		banded iron formations (Fe and Mn)	
b) carbonates		ferroan carbonates, dolostones, limestones	
c) sulfates		barite, gypsum	
d) evaporites		halides, borates	
e) carbonate skarns		diopside, tremolite, grossularite, wollastonite, scapolite	
		iated mineralogy (>2.5 Ga to present)	
7. Paleoproterozoic atmospheric changes	2.5 to 1.9 Ga	>2000 new oxide/hydroxide species, especially ore minerals	>4000
surface oxidation			
8. Intermediate ocean	1.9 to 1.0 Ga	minimal mineralogical innovation	>4000
9. Neoproterozoic biogeochemical changes	1.0 to 0.542 Ga		>4000
a) glaciation		extensive ice deposition, but few new minerals	
b) post-glacial oxidation		extensive oxidative weathering of all surface rocks	
10. Phanerozoic Era	0.542 Ga to present		4300+
a) biomineralization		extensive skeletal biomineralization of calcite, aragonite,	
		dolomite, hydroxylapatite, and opal	
b) bio-weathering		increased production of clay minerals, soils	

5. Importance of water for geomaterials

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Water:



important in the *formation* of many geomaterials ↓

→ important for the *control* of these geomaterials

Lime

Slaking of burned lime (CaO) with water → hydrated lime Ca(OH)2

Gypsum

Removal water of crystallisation → burned gypsum (plaster) → rehydrate dry plaster by adding water: CaSO₄·2H₂O + heat → CaSO₄·0.5H₂O + 1.5H₂O

Clay

Adding water: processable via plastic ductility (swellable clay) Burning: resilient ceramic.

6. Coevolution of geo- and biosphere driven by water

Minerals and the Origin of Life



Life- mediated Origin of Minerals



6. Coevolution of geo- and biosphere

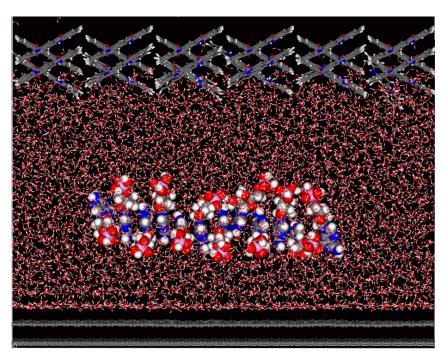
6.1. Minerals and the Origin of Life

- Mineral surfaces:
 - concentration of organic molecules (e.g. clay minerals) [dx.doi.org/10.5772/50172]
 - selection (chirality) of L-amino acids (e.g. calcite) [doi.org/10.1073/pnas.101085998]
 - stabilyzing amino acids (e.g. pyrite) [doi.org/10.1021/ja063295a]
 - catalyzing (bio)chemical reactions (e.g. pyrite) [doi.org/10.1038/344387a0]
- Mineral pores: prebiotic reaction vessels (e.g. pumice) [doi.org/10.1089/ast.2010.0546]
- Mineral nanoconfinements: biopolymerization within aqueous particle suspensions (e.g. graphite, magnetite, silica)

Mineral nanoconfinements:

→ biopolymerization within aqueous particle suspensions

(e.g. graphite, PHAs, magnetite, silica)



6.2. Life-mediated Origin of Minerals

- Great Oxidation Event (~ 2,5 Ga)
 After rise of oxygenic photosynthesis: no rise of atmosphereic O2
 → oxidation of dissolved Fe and other oxidizable compounds (e.g. Methane gets oxidized to CO2 (much weaker greenhouse gas) and H2O → Huronian Glaciation) which are constantly resupplied via volcanism + erosion
 - → numerous new mineral species and morphologies as oxidiation products (oxides, hydroxides)
- <u>Phanerozoic biomineralization (since ~ 0,54 Ga)</u>
 "sudden" rise of most prevalent biominerals (calcite, dolomite, aragonite, apatite).

Limestone Alps: Part of geosphere, but material originated from biosphere. Biomineralization bound CO₂.

Water + minerals → origin of life Oxide minerals / biominerals ← water + life →



Summary

Water as a Meta-Geomaterial:

- → Origin on mineral particles
- → Transport to Earth via chondrites
- → Remains on Earth due to Fe, U
- → Boosts evolution of mineral species and morphologies
- → Drives geobiology
- → Enables to control properties of geomaterials due to its role in their origin